

Review of Hydraulic Properties for Sediments in the 200 Areas

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Presentation Topics ...

- 200 Area database on soil properties ...
- Variability and scaling for 200 Area samples ...
- Recent data at the ILAW site in 200 East Area ...
- Anisotropy results for sediments at the ILAW site

200 Areas Database ...

- 183 samples from 12 sites in 200 Areas have data on PSD, bulk density, moisture retention and saturated K
- Moisture retention and saturated K data corrected for gravel content
- Data grouped into six categories: SS, sand mixed w/ finer fraction; S, sand; SSG, sand and gravel mixed w/ finer fraction; GS, gravelly sand; SG1, sandy gravel w/ gravel fraction < 60%; SG2, sandy gravel w/ gravel fraction >60%
- Variability and scaling results available for the 183 samples
- Data on PSD, bulk density, moisture retention, saturated and unsaturated K available for 42 samples

Location of Sites in 200 Areas Plateau

Correcting Saturated K for Gravels: Bouwer and Rice (1984)

$$K_b = K_s \left(\frac{e_b}{e_f} \right)$$

K_b = saturated K for the bulk sample

K_s = saturated K for the < 2 mm fraction

e_b = void ratio for the bulk sample

e_f = void ratio for the < 2 mm fraction

Correcting Moisture Retention for Gravels: Gardner (1986)

$$q_b = \frac{W_f r_b / r_w}{1 + m_g / m_f}$$

\hat{e}_b =volumetric moisture content of the bulk soil including gravel

w_f =laboratory-measured gravimetric moisture content of the fine fraction (< 2 mm)

\tilde{n}_b =bulk density of the entire sample (including gravel)

\tilde{n}_w =density of water

m_f and m_g =respective dry masses of the fines and gravel as recovered from the bulk field sample

Scaling Relationships: Vogel et al. (1991)

$$K(h) = \mathbf{a}_k K^*(h^*)$$

$$\mathbf{q}(h) = \mathbf{q}_r + \mathbf{a}_q [\mathbf{q}^*(h^*) - \mathbf{q}_r^*]$$

$$h = \mathbf{a}_h h^*$$

α_e , α_h , α_k =mutually independent scaling factors for moisture content, pressure head and hydraulic conductivity, respectively.

For texturally similar soils, scaling factors are:

$$\mathbf{a}_k = K_s / K_s^*$$

$$\mathbf{a}_q = (\mathbf{q}_s - \mathbf{q}_r) / (\mathbf{q}_s^* - \mathbf{q}_r^*)$$

$$\mathbf{a}_h = h_c / h_c^*$$

Macroscopic Anisotropy Relationships: Polmann (1990)

$$\langle \ln K \rangle = \langle \ln K_s \rangle - A \langle \phi \rangle - \mathbf{S}_{\ln K_s}^2 \mathbf{I} [p - p^2 \langle \phi \rangle - \mathbf{z}^2 \langle \phi \rangle] / (1 + A \mathbf{I})$$

$$\mathbf{S}_{\ln K}^2 = \mathbf{S}_{\ln K_s}^2 [(1 - p \langle \phi \rangle)^2 + \mathbf{z}^2 \langle \phi \rangle^2] / (1 + A \mathbf{I})$$

$$K_h^{eq} = \exp[\langle \ln K \rangle + (\mathbf{S}_{\ln K}^2 / 2)]$$

$$K_v^{eq} = \exp[\langle \ln K \rangle - (\mathbf{S}_{\ln K}^2 / 2)]$$

where $\mathbf{S}_{\ln K}^2$ = variance of log unsaturated conductivity (which depends on mean tension),

$\langle \phi \rangle$ = mean tension,

$\mathbf{S}_{\ln K_s}^2$ = variance of $\ln K_s$

$\langle \ln K_s \rangle$ = mean of $\ln K_s$,

p = slope of the $\hat{\alpha}$ versus $\ln K_s$ regression line,

$\hat{\alpha} = \bar{\alpha} / \bar{\sigma}_{\ln K_s}$,

$\bar{\sigma}_{\hat{\alpha}}$ = standard deviation of the residuals in the $\hat{\alpha}$ versus $\ln K_s$ regression,

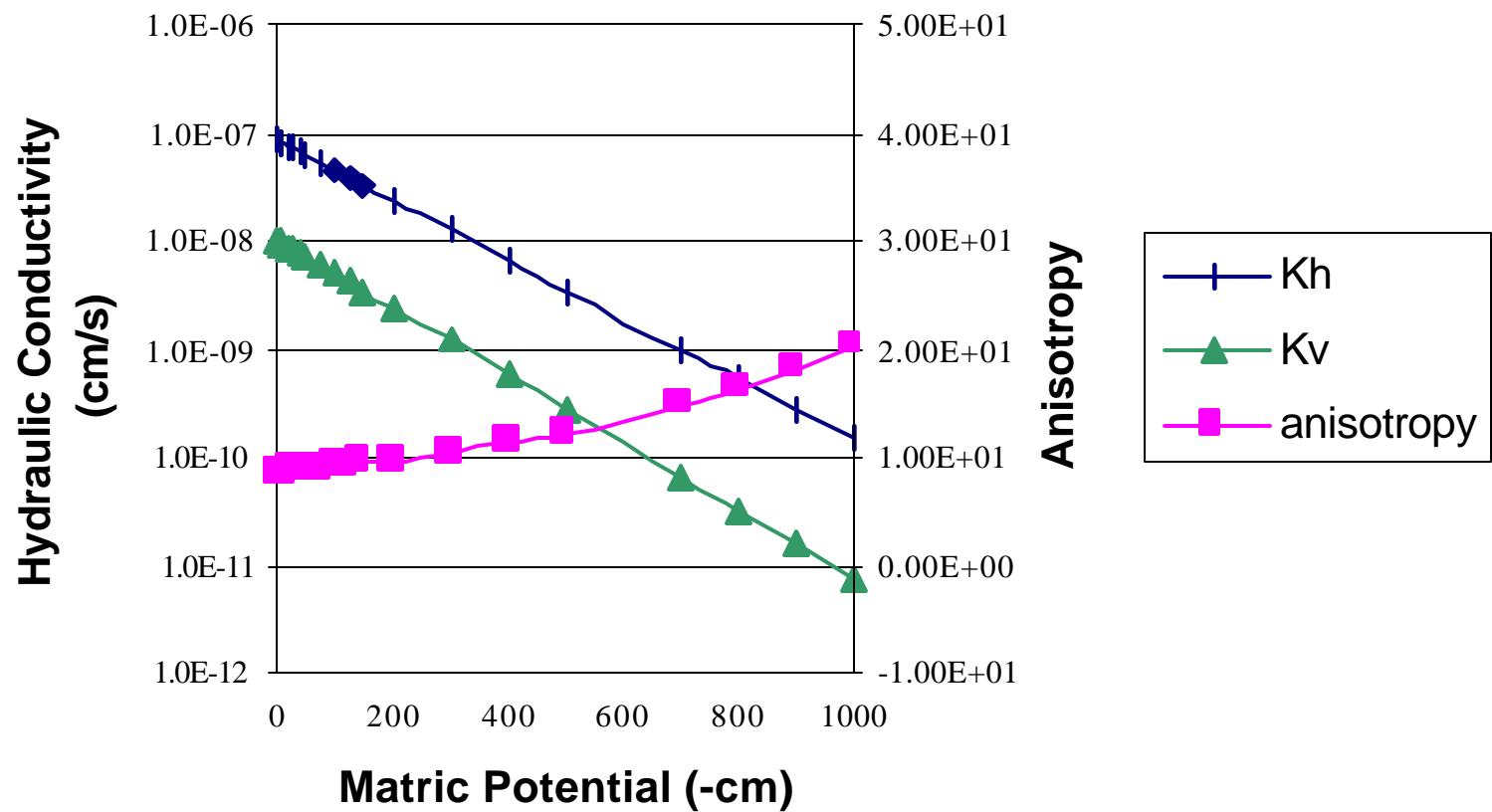
A = mean slope, $\hat{\alpha}$, for $\ln K_s$ vs. ϕ ,

\hat{e} = vertical correlation lengths for $\ln K_s$ (assumed to be same as that of $\hat{\alpha}$),

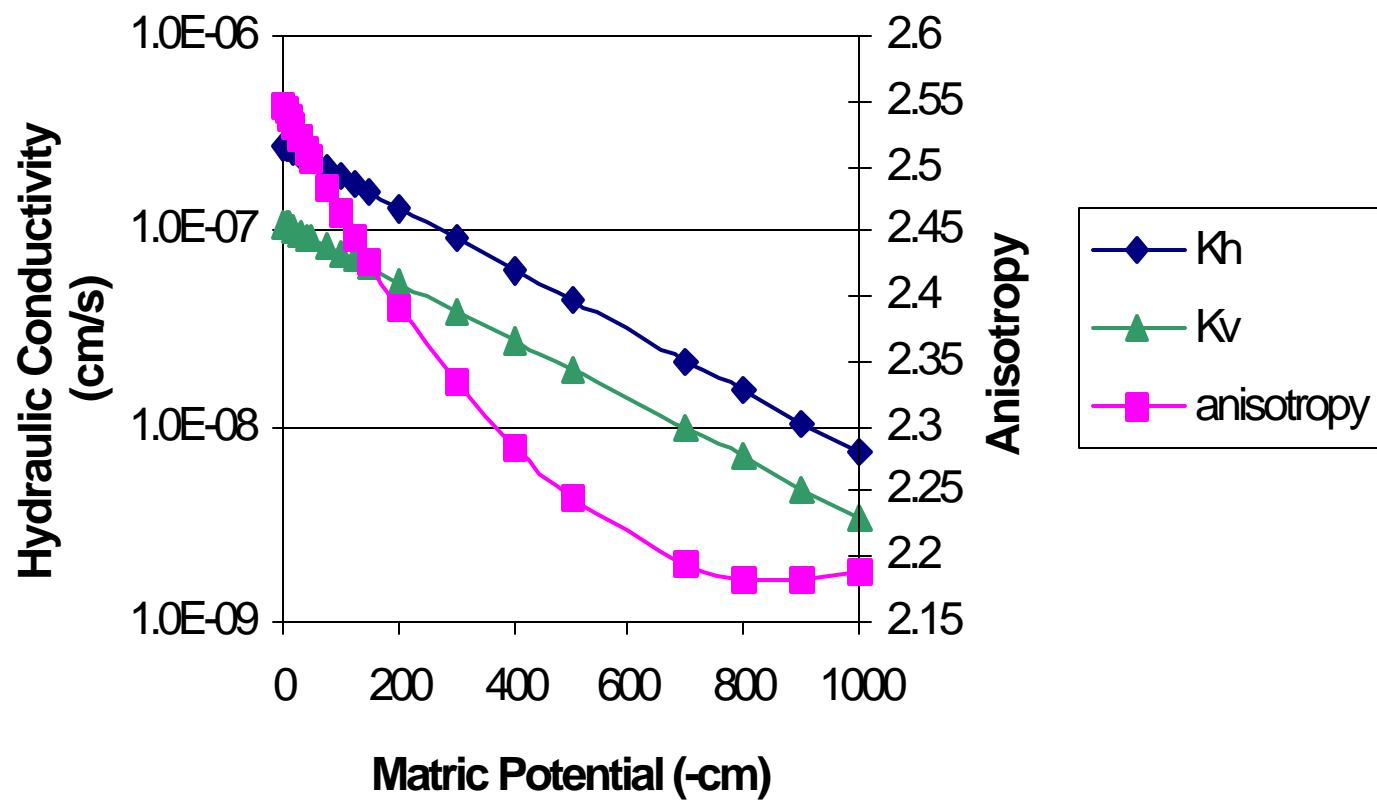
K_h^{eq} = equivalent unsaturated horizontal conductivity, and

K_v^{eq} = equivalent unsaturated vertical conductivity.

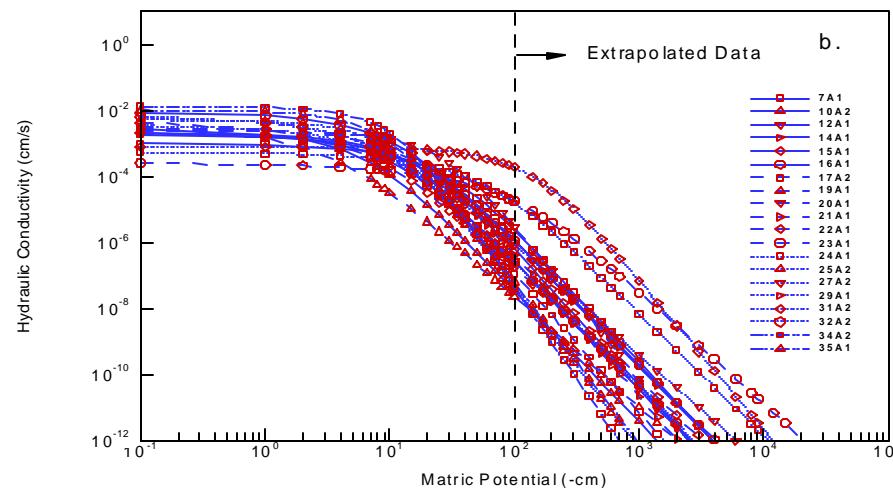
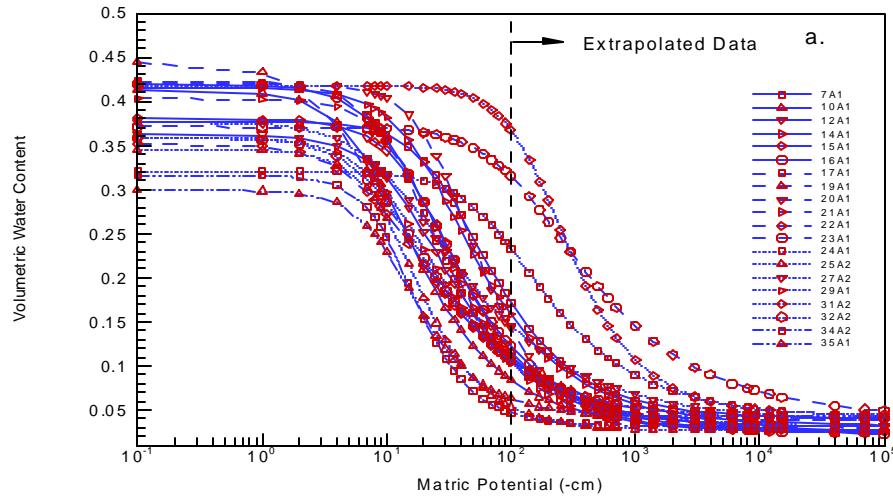
Macroscopic Anisotropy for the Hanford Sandy Sequence at the ILAW Site in 200 E Area



Macroscopic Anisotropy for the Hanford Gravelly Sequence at the ILAW Site in 200 E Area



Moisture Retention and Unsat K Data for the Hanford Sandy Sequence at the ILAW Site



Moisture Retention and Unsat K Data for the Gravelly Sequence at the ILAW Site

